Partial Differential Equations

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Functions of More Than One Variable

 Almost all processes relevant in geosciences are described by variables varying in time and space. The spatial component is a scalar in case of a one-dimensional description and a vector in case of a two- or three-dimensional description.

Examples:

 $T(\vec{x}, t)$ as the temperature

- $p(\vec{x}, t)$ as the fluid pressure in a reservoir
- $\rho(\vec{x}, t)$ as the density in a gas
- $\vec{v}(\vec{x},t)$ as the flow velocity in a fluid
- Spatial interactions often refer to the spatial variation in the variables. Examples:

Heat conduction is driven by spatial differences in temperature. Fluid flow is driven by spatial differences in pressure.



Partial Derivatives

- If a function u depends on more than one variable, e.g., u(x₁, x₂, x₃, t) (or shorter u(x, t)), the derivative with respect to one of the variables (while the others are constant) is called partial derivative.
- Partial derivatives are written with the symbol ∂ , e.g.,

$$\frac{\partial}{\partial x_1}u(\vec{x},t), \ \frac{\partial}{\partial x_2}u(\vec{x},t), \ \frac{\partial}{\partial x_3}u(\vec{x},t), \ \text{and} \ \frac{\partial}{\partial t}u(\vec{x},t).$$

• Partial derivatives are computed by assuming that the other variables are constant.



Examples of Partial Derivatives

• Density of an ideal gas

$$\rho(p, T) = \frac{M}{R} \frac{p}{T} \quad \text{with} \quad \begin{array}{l} M = \text{molar mass} \\ R = \text{gas constant} \\ \\ \frac{\partial}{\partial p} \rho(p, T) = \\ \end{array}, \quad \begin{array}{l} \frac{\partial}{\partial T} \rho(p, T) = \\ \end{array}$$

• One-dimensional harmonic wave

$$u(x, t) = A \sin(\omega t - kx)$$
 with $\begin{aligned} A &= & \text{amplitude} \\ \omega &= & \text{angular frequency} \\ k &= & \text{wave number} \end{aligned}$

$$\frac{\partial}{\partial x}u(x,t) = , \quad \frac{\partial}{\partial t}u(x,t) =$$

The Gradient

The partial derivatives with respect to the spatial coordinates are often subsumed in a vector

$$\operatorname{grad} u(\vec{x}, t) = \nabla u(\vec{x}, t) = \begin{pmatrix} \frac{\partial}{\partial x_1} u(\vec{x}, t) \\ \frac{\partial}{\partial x_2} u(\vec{x}, t) \\ \frac{\partial}{\partial x_3} u(\vec{x}, t) \end{pmatrix}$$

Examples:





Properties of the Gradient

- $\nabla u(\vec{x})$ is normal to the lines (in 2D) or the surfaces (in 3D) where $u(\vec{x})$ is constant.
- $\nabla u(\vec{x})$ points in direction of steepest increase of $u(\vec{x})$.
- The length of $\nabla u(\vec{x})$ is the slope of $u(\vec{x})$ in direction of steepest increase.

What Is a Partial Differential Equation?

A partial differential equation (PDE) is an equation for an unknown function depending on more than variable involving partial derivatives.

A differential equation for a function of only one variable is called ordinary differential equation (ODE).

Examples of Partial Differential Equations



The One-Dimensional Advection Equation

$$\frac{\partial}{\partial t}u(x,t) = -v \frac{\partial}{\partial x}u(x,t)$$

The Equation of Continuity (mass conservation) in a Fluid

$$\begin{aligned} \frac{\partial}{\partial t}\rho(\vec{x},t) &= -\operatorname{div}\left(\rho(\vec{x},t)\,\vec{v}(\vec{x},t)\right) \\ &= -\frac{\partial}{\partial x_1}\left(\rho(\vec{x},t)\,v_1(\vec{x},t)\right) - \frac{\partial}{\partial x_2}\left(\rho(\vec{x},t)\,v_2(\vec{x},t)\right) \\ &- \frac{\partial}{\partial x_3}\left(\rho(\vec{x},t)\,v_3(\vec{x},t)\right) \end{aligned}$$



The Heat Conduction Equation

Simplest version (1D with constant parameters):

$$\frac{\partial}{\partial t}T(x,t) = \kappa \frac{\partial^2}{\partial x^2}T(x,t)$$

General version (3D):

$$\rho c \frac{\partial}{\partial t} T(\vec{x}, t) = \operatorname{div} \left(\lambda \nabla T(\vec{x}, t) \right)$$
$$= \frac{\partial}{\partial x_1} \left(\lambda \frac{\partial}{\partial x_1} T(\vec{x}, t) \right) + \frac{\partial}{\partial x_2} \left(\lambda \frac{\partial}{\partial x_2} T(\vec{x}, t) \right)$$
$$+ \frac{\partial}{\partial x_3} \left(\lambda \frac{\partial}{\partial x_3} T(\vec{x}, t) \right)$$



The Navier-Stokes Equations of a Viscous Fluid (without gravity)

$$\rho\left(\frac{\partial}{\partial t}\vec{v}(\vec{x},t) + (\vec{v}(\vec{x},t)\cdot\nabla)\vec{v}(\vec{x},t)\right) = -\nabla p(\vec{x},t) + \eta\Delta\vec{v}(\vec{x},t)$$

with

$$\vec{v}(\vec{x},t) \cdot \nabla \vec{v}(\vec{x},t) = v_1(\vec{x},t) \frac{\partial}{\partial x_1} \vec{v}(\vec{x},t) + v_2(\vec{x},t) \frac{\partial}{\partial x_2} \vec{v}(\vec{x},t) + v_3(\vec{x},t) \frac{\partial}{\partial x_3} \vec{v}(\vec{x},t)$$
$$+ v_3(\vec{x},t) \frac{\partial}{\partial x_3} \vec{v}(\vec{x},t)$$
$$\Delta \vec{v}(\vec{x},t) = \frac{\partial^2}{\partial x_1^2} \vec{v}(\vec{x},t) + \frac{\partial^2}{\partial x_2^2} \vec{v}(\vec{x},t) + \frac{\partial^2}{\partial x_3^2} \vec{v}(\vec{x},t)$$



Treatment of the Time Coordinate

The procedure is basically the same as for the time in ordinary differential equations:

- If second-order derivatives occur, the first-order derivatives must be introduced as separate variables.
- Approximate solutions are computed step by step (in steps of length δt), starting from the initial time t_0 .
- The time derivative is approximated by

$$\frac{\partial}{\partial t}u(\vec{x},t) \approx \frac{u(\vec{x},t+\delta t)-u(\vec{x},t)}{\delta t}$$



Treatment of the Time Coordinate

- All schemes (explicit and fully implicit Euler, mixed, e.g., Crank-Nicholson) can be used.
- If the explicit Euler scheme is used, the solution u(x, t + δt) can be directly obtained from u(x, t) by computing partial derivatives of u(x, t) with respect to the spatial coordinates x.
- For all implicit schemes, a partial differential equation with respect to the spatial coordinates \vec{x} remains to be solved in each timestep.



The Finite-Difference Method in One Dimension

One-dimensional case: only one spatial coordinate x (and time)

First step:

- Discrete points x₁, x₂, ..., x_n are defined on the considered part of the x-axis (from the left-hand boundary to the right-hand boundary).
- These points are called nodes and are the points where an approximate solution will be computed.
- The nodes may be equidistant (having all the same distance δx) or not.



The Finite-Difference Method in One Dimension

Second step: The partial derivatives $\frac{\partial}{\partial x}u(x, t)$ are approximated by difference quotients.

Right-handed difference quotient:

$$\frac{\partial}{\partial x}u(x,t) \approx \frac{u(x+\delta x,t)-u(x,t)}{\delta x}$$

Left-handed difference quotient:

$$\frac{\partial}{\partial x}u(x,t) \approx \frac{u(x,t)-u(x-\delta x,t)}{\delta x}$$

Numerics of Partial Differential Equations



The Finite-Difference Method in One Dimension

Central (symmetric) difference quotient:

$$\frac{\partial}{\partial x}u(x,t) \approx \frac{u(x+\delta x,t)-u(x-\delta x,t)}{2\delta x}$$

or

$$\frac{\partial}{\partial x}u(x,t) \approx \frac{u(x+\frac{\delta x}{2},t)-u(x-\frac{\delta x}{2},t)}{\delta x}$$

or

$$\frac{\partial}{\partial x}u(x+\frac{\delta x}{2},t) \approx \frac{u(x+\delta x,t)-u(x,t)}{\delta x}$$

or

$$\frac{\partial}{\partial x}u(x-\frac{\delta x}{2},t) \approx \frac{u(x,t)-u(x-\delta x,t)}{\delta x}$$

Only the first version can be applied directly, the others are only useful for combining them to second-order derivatives.



The Finite-Difference Method in One Dimension

- The accuracy of all these difference quotients decreases with increasing δx .
- As long as there is no preferred direction, right-hand and left-hand difference quotients are equivalent.
- Central difference quotients provide a higher accuracy than the one sided versions.

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General Concept

Consider the amount of anything that is worth keeping track of (mass, energy, \ldots) in a volume

rate of change of the amount stored within the volume

amount per time entering at the boundaries amount per time leaving at the boundaries

+

amount per time produced within the volume

amount per time removed within the volume



Density and Flux Density

Density $u(\vec{x}, t) = \text{amount per volume}$

Flux density $\vec{q}(\vec{x}, t)$ = amount passing a surface per time and surface area

 \vec{q} is a vector, so that the amount per time passing a (small) surface of size A with a unit normal vector \vec{n} is $\vec{q} \cdot \vec{n}A$.



The Equation of Continuity

Balance of a the amount contained in a cube:



where Q is the rate of production (amount per time and volume)



The General Balance Equation

$$\frac{\partial}{\partial t}u(\vec{x},t) = -\operatorname{div}\vec{q}(\vec{x},t) + Q(\vec{x},t)$$

where

$$\operatorname{div} \vec{q}(\vec{x}, t) = \frac{\partial}{\partial x_1} q_1(\vec{x}, t) + \frac{\partial}{\partial x_2} q_2(\vec{x}, t) + \frac{\partial}{\partial x_3} q_3(\vec{x}, t)$$

= divergence of the flux density \vec{q}



The Equation of Advection

Assume that the considered amount moves at a given velocity $\vec{v}(\vec{x}, t)$:

$$\vec{q}(\vec{x},t) = u(\vec{x},t) \vec{v}(\vec{x},t)$$

$$\frac{\partial}{\partial t}u(\vec{x},t) = -\operatorname{div}\left(u(\vec{x},t)\,\vec{v}(\vec{x},t)\right) + Q(\vec{x},t)$$

In 1D:

$$\frac{\partial}{\partial t}u(x,t) = -\frac{\partial}{\partial x}(u(x,t)\vec{v}(x,t)) + Q(x,t)$$

The equation of advection is of first order in both time and space (hyperbolic differential equation).



The Equation of Advection

Simplest version: $\vec{v}(\vec{x}, t) = \text{const.}, Q(\vec{x}, t) = 0$:

$$\frac{\partial}{\partial t}u(\vec{x},t) = -\vec{v}\cdot\nabla u(\vec{x},t)$$

In 1D:

$$\frac{\partial}{\partial t}u(x,t) = -v\frac{\partial}{\partial x}u(x,t)$$



The Diffusion Equation

- Flux follows the direction of steepest descent of the density $u(\vec{x}, t)$.
- 2 The flux density is proportional to the decrease of density per length.

$$ec{q}(ec{x},t) = -D
abla u(ec{x},t) = -D\left(egin{array}{c} rac{\partial}{\partial x_1}u(ec{x},t)\\ rac{\partial}{\partial x_2}u(ec{x},t)\\ rac{\partial}{\partial x_3}u(ec{x},t) \end{array}
ight)$$

with

 $D = \text{diffusivity (coefficient of diffusion) } \left[\frac{m^2}{s}\right]$



The Diffusion Equation

Insert flux density into the balance equation:

$$\frac{\partial}{\partial t}u(\vec{x},t) = -\operatorname{div}\vec{q}(\vec{x},t) + Q(\vec{x},t)$$
$$= \operatorname{div}(D\nabla u(\vec{x},t)) + Q(\vec{x},t)$$

In 1D:

$$\frac{\partial}{\partial t}u(x,t) = \frac{\partial}{\partial x}\left(D\frac{\partial}{\partial x}u(x,t)\right) + Q(x,t)$$

The diffusion equation is of first order in time and of second order in space (parabolic differential equation).



The Diffusion Equation

Simplest version: D = const., $Q(\vec{x}, t) = 0$:

$$\begin{aligned} \frac{\partial}{\partial t}u(\vec{x},t) &= D\operatorname{div}\nabla u(\vec{x},t) \\ &= D\left(\frac{\partial^2}{\partial x_1^2}u(\vec{x},t) + \frac{\partial^2}{\partial x_2^2}u(\vec{x},t) + \frac{\partial^2}{\partial x_3^2}u(\vec{x},t)\right) \\ &= D\Delta u(\vec{x},t) \end{aligned}$$

where $\Delta = \frac{\partial^2}{\partial x_1^2} + \frac{\partial^2}{\partial x_2^2} + \frac{\partial^2}{\partial x_3^2}$ is the Laplace operator. In 1D:

$$\frac{\partial}{\partial t}u(\vec{x},t) = D\frac{\partial^2}{\partial x^2}u(x,t)$$



Initial Conditions and Boundary Conditions

- Time is distinct from the spatial coordinates as it is directed.
- For a unique solution, partial differential equations require Initial conditions: The solution for all points x of the domain at a time t₀ must be given.
 Boundary conditions: The solution (or, e.g., its derivatives) must be given at (least at a part of) the boundary of the domain for all times t > t₀.

Numerics of Partial Differential Equations



The Finite-Difference Method in Two Dimensions

Two-dimensional case: Spatial coordinates x_1 and x_2 .

First step: The domain is covered by a lattice where the lines are parallel to the coordinate axes.







The Finite-Difference Method in Two Dimensions

Second step: The partial derivatives $\frac{\partial}{\partial x_1}u(x, t)$ and $\frac{\partial}{\partial x_2}u(x, t)$ are approximated by difference quotients.

Right-handed difference quotients:

$$\frac{\partial}{\partial x_1}u(x_1, x_2, t) \approx \frac{u(x_1 + \delta x, x_2, t) - u(x_1, x_2, t)}{\delta x}$$
$$\frac{\partial}{\partial x_2}u(x_1, x_2, t) \approx \frac{u(x_1, x_2 + \delta x, t) - u(x_1, x_2, t)}{\delta x}$$

Left-handed difference quotients:

$$\frac{\partial}{\partial x_1} u(x_1, x_2, t) \approx \frac{u(x_1, x_2, t) - u(x_1 - \delta x, x_2, t)}{\delta x}$$
$$\frac{\partial}{\partial x_2} u(x_1, x_2, t) \approx \frac{u(x_1, x_2, t) - u(x_1, x_2 - \delta x, t)}{\delta x}$$



The Finite-Difference Method in Two Dimensions

Central difference quotients:

$$\begin{aligned} \frac{\partial}{\partial x_1} u(x_1, x_2, t) &\approx \quad \frac{u(x_1 + \delta x, x_2, t) - u(x_1 - \delta x, x_2, t)}{2\delta x} \\ \frac{\partial}{\partial x_1} u(x_1, x_2, t) &\approx \quad \frac{u(x_1 + \frac{\delta x}{2}, x_2, t) - u(x_1 - \frac{\delta x}{2}, x_2, t)}{\delta x} \\ \frac{\partial}{\partial x_2} u(x_1, x_2, t) &\approx \quad \frac{u(x_1, x_2 + \delta x, t) - u(x_1, x_2 - \delta x, t)}{2\delta x} \\ \frac{\partial}{\partial x_2} u(x_1, x_2, t) &\approx \quad \frac{u(x_1, x_2 + \frac{\delta x}{2}, t) - u(x_1, x_2 - \frac{\delta x}{2}, t)}{\delta x} \end{aligned}$$

and all other variants discussed in the one-dimensional case



The Energy Balance

$$\frac{\partial}{\partial t}\epsilon(\vec{x},t) = -\operatorname{div}\vec{q}(\vec{x},t) + Q(\vec{x},t)$$

where

$$\begin{aligned} \epsilon(\vec{x}, t) &= \text{ energy density } \left[\frac{J}{m^3}\right] \\ \vec{q}(\vec{x}, t) &= \text{ heat flux density } \left[\frac{W}{m^2}\right] \\ Q(\vec{x}, t) &= \text{ rate of production of thermal energy per volume } \left[\frac{W}{m^3}\right] \end{aligned}$$

The Heat Equation



The Specific Heat Capacity

The specific heat capacity

$$c = \frac{1}{\rho} \frac{\partial \epsilon}{\partial T}$$

describes the change in thermal energy with temperature.

$$\frac{\partial}{\partial t}\epsilon(\vec{x},t) = \rho c \frac{\partial}{\partial t}T(\vec{x},t) = -\operatorname{div}\vec{q}(\vec{x},t) + Q(\vec{x},t)$$
Water: $c = 4180 \frac{J}{\text{kgK}}$
Rocks: $c = 800-1000 \frac{J}{\text{kgK}}$



The Three Mechanisms of Heat Transport

Heat conduction: Driven by spatial differences in temperature. Advective transport: Heat is carried by a moving (flowing) medium.

$$ec{q}(ec{x},t) \;=\; \epsilon(ec{x},t)\,ec{v} \;=\;
ho c \; T(ec{x},t)\,ec{v}$$

Radiation: Does not require a medium, but is only significant at very high temperatures.



Fourier's Law of Heat Conduction (1822)

- Heat flux follows the direction of steepest descent of the temperature field T(x, t).
- The heat flux density is proportional to the decrease of temperature per length.

$$\vec{q}(\vec{x},t) = -\lambda \nabla T(\vec{x},t) = -\lambda \left(\begin{array}{c} \frac{\partial}{\partial x_1} T(\vec{x},t) \\ \frac{\partial}{\partial x_2} T(\vec{x},t) \\ \frac{\partial}{\partial x_3} T(\vec{x},t) \end{array} \right)$$

with

$$\lambda$$
 = thermal conductivity $\left[\frac{W}{mK}\right]$



Typical Values of λ

Material	$\lambda \left[\frac{W}{m K}\right]$
diamond	2300
iron	80
quartz	1.4
sand	0.6
expanded polystyrene	0.033
water	0.6
air	0.026

Rocks	$\lambda \left[\frac{W}{m K}\right]$
granite	2.8
basalt	2
dolomite	2.5
limestone	2.5
sandstone	2.5
shale	2
widely used value	2.5



The Full Heat Equation (conduction, advection, production)

Energy balance + Fourier's law + advective heat flux

$$\oint \rho c \frac{\partial}{\partial t} T(\vec{x}, t) = \operatorname{div} \left(\lambda \nabla T(\vec{x}, t) - \rho c T(\vec{x}, t) \vec{v} \right) + Q$$



The Thermal Diffusivity

If ρ , c, and λ are constant (mass balance implies div $\vec{v} = 0$ then), the heat equation can be simplified to

$$\begin{aligned} \frac{\partial}{\partial t}T(\vec{x},t) &= \kappa \Delta T(\vec{x},t) - \vec{v} \cdot \nabla T(\vec{x},t) + \frac{Q}{\rho c} \\ &= \kappa \left(\frac{\partial^2}{\partial x_1^2}T(\vec{x},t) + \frac{\partial^2}{\partial x_2^2}T(\vec{x},t) + \frac{\partial^2}{\partial x_3^2}T(\vec{x},t)\right) \\ &- \vec{v} \cdot \nabla T(\vec{x},t) + \frac{Q}{\rho c} \end{aligned}$$

with the thermal diffusivity

$$\kappa = \frac{\lambda}{\rho c}$$

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Water:
$$\kappa = 1.4 \times 10^{-7} \frac{\text{m}^2}{\text{s}}$$

Rocks: $\kappa \approx 10^{-6} \frac{\text{m}^2}{\text{s}} \approx 30 \frac{\text{m}^2}{\text{a}}$



Boundary Conditions

Boundary conditions for a second-order (concerning space) differential equation concern the values or their derivatives (or a combination of both). Dirichlet boundary conditions define the temperatures at the boundaries:

 $T(\vec{x},t) = T_b$

Interpretation: System is coupled to a big reservoir held a a given temperature.

The Heat Equation



Boundary Conditions

(Von) Neumann boundary conditions define the conductive heat flux density across the boundaries:

$$-\lambda \nabla T(\vec{x},t) \cdot \vec{n} = q_b$$

where \vec{n} is the outer unit normal vector on the surface. In 1D:

$$\pm \lambda \frac{\partial}{\partial x} T(x, t) = q_b$$

Homogeneous (von) Neumann boundary condition: $q_b = 0$

$$\nabla T(\vec{x},t) \cdot \vec{n} = 0$$
 (2D/3D), $\frac{\partial}{\partial x}T(x,t) = 0$ (1D)



Boundary Conditions

Mixed boundary conditions define a combination of temperature and temperature gradient normal to the surface.

Examples:

• Total (conductive + advective) heat lux density across the boundary is given:

$$(-\lambda \nabla T(\vec{x}, t) + \rho c T(\vec{x}, t) \vec{v}) \cdot \vec{n} = q_b$$

• Radiating surface (without advection):

$$-\lambda \nabla T(\vec{x},t) \cdot \vec{n} = \sigma T(\vec{x},t)^4$$