Hydrogeology Ground Penetrating Radar

Jakob Wilk

Institut für Geo- und Umweltnaturwissenschaften Albert-Ludwigs-Universität Freiburg



Basics

- GPR is a geophysical measurement technique which can be applied to explore near-surface underground structures.
- The measurement principle is based on the transmittance (transmitting antenna) of high-frequency (50MHz – 2.5 GHz) electromagnetic pulses into the ground. Reflection or Refraction at inhomogeneities are received (receiving antenna).
- Travel time and amplitude give information about the structure and depth of the inhomogeneities.

Applications

- Geological
 - Detection of cavities and fissures
 - •Mapping of superficial deposits, subsidence, soil stratigraphy, fractures, etc.
 - •Mineral exploration and resource evaluation
 - •(Paleo-) Lake and riverbed mapping
 - Environmental
 - Contaminant plume mapping
 - Location of gas leaks
 - Groundwater-level
 - Mapping of pollutants within groundwater
 - Glaciological
 - Ice thickness and snow stratigraphy

Applications



Radio Sounding Skelton Glacier 1964

Applications



Applications

Engineering and construction

- Void detection
- Road pavement
- Concrete testing
- Location of public utilities
- Archaeology
 - Location of buried structures
 - •Pre-excavation mapping
 - •Location of graves, crypts, etc.
- Forensic science
 - •Location of buried corpses





Interferogram produced using <u>ERS-2</u> data from 13 August and 17 September 1999, spanning the 17 August <u>Izmit</u> (Turkey) earthquake. (NASA/JPL-Caltech)

Shallow radar sounder - SHARAD



Radargram of north pole layered deposits (ASI, Mars Reconnaissance Orbiter)



Propagation of electromagnetic waves

The emitted electromagnetic waves are affected by

- •Dielectric properties ϵ_r
- •Magnetic susceptibility μ_r
- •Electrical conductivity σ_e

Reflection coefficient

inclination is:

•Describes how much of an electromagnetic wave is reflected by an impedance discontinuity in the transmission medium.

•Of a plane (fraction of the reflected energy at the plane for normal

$$\mathbf{R} = \frac{v_1 - v_2}{v_1 + v_2} = \frac{\sqrt{\varepsilon_2} - \sqrt{\varepsilon_1}}{\sqrt{\varepsilon_2} + \sqrt{\varepsilon_1}}$$

Propagation of electromagnetic waves

Material	ε _r	Material	ε _r
Air	1	Saturated Sand	25
Ice	3	Silts	30
Dry Sand	5	Clays	40
Granite	6	Fresh Water	80
Dry Salt	6	Sea Water	80
Limestone	8	Metal	~
Shale	15	Table from Telford 1990 Applied Geophysics	

Propagation of electromagnetic waves



Measurement principle

The standard system consist of a transmitting and receiving antenna. The radar wave
propagates through the soil while the velocity of the wave depends on the dielectric
properties of the ground. At interfaces, where the dielectric properties of the
different media change, the electromagnetic wave is partially reflected. The travel
time and amplitude of the wave is recorded by the receiving antenna.

Measurement principle



Measurement principle

- In a standard GPR measurement, the antennas are pulled along the survey track while traces are triggered at a fixed interval by a measurement wheel which is connected to the back of the antenna. This results in a series of traces which are finally displayed by the measurement software as a function of position and time in a so-called radargram.
- The sampling interval and the antenna frequency has to follow the Nyquist-Shannon sampling theorem:

•
$$f_{Nyq} = \frac{1}{2\Delta t} > fmax (f_{max} is \sim antenna frequency + 50\%)$$



Measurement principle

• The signals of the emitted waves reach the receiving antenna at different times. Plotting the recorded amplitudes as a function of time results in a so-called "trace".



Measurement principle



Common offset (CO)

- Is the simplest and most widespread GPR measurement technique. Transmitter and receiver antenna are moved along the survey track while the distance between both antenna is kept constant. Electromagnetic pulses are emitted at equidistant intervals which are controlled by the survey wheel.
- From the measured travel time *t* of the reflected signal the depth *d* of a horizontal reflector can be determined by:

$$s = 2 \cdot \sqrt{d^2 + \left(\frac{a}{2}\right)^2} = \sqrt{4d^2 + a^2}$$

with a as distance between antennas.

• Assuming a homogeneous medium the travel time is determined by:

$$t = \frac{s}{v} = \frac{\sqrt{4d^2 + a^2}}{\frac{c}{\sqrt{\varepsilon}}}$$





Multi offset: Common midpoint

 In a common-midpoint measurement (CMP) transmitter and receiver are moved away from each other in equidistant steps. The resulting radargram displays the travel time as a function of the antenna separation.

Multi offset: Common midpoint

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Multi offset: Common midpoint

- Air and ground wave travel directly between the antenna.
- Linear relationship between the travel time t and the antenna separation a with the constant of proportionality $\frac{1}{v}$:

$$t = \frac{a}{v}$$
, with $v = c$ (airwave) and $v = \frac{c}{\sqrt{\varepsilon}}$ (groundwave)

 Reflector depth below the midpoint between the transmitting and the receiving antenna: the relation between travel time t reflector depth is given by:

$$s = 2 \cdot \sqrt{d^2 + \left(\frac{a}{2}\right)^2} = \sqrt{4d^2 + a^2}$$

• Plotting the measured data in a t^2 -a²-diagram, leads to a linear relationship between t and a: 1 $4h^2$

$$t^2 = \frac{1}{v^2}a^2 + \frac{4h^2}{v^2}$$

Multi offset: Common midpoint





Multi offset: Common midpoint



Multi offset: Wide Angle Reflection and Refraction (WARR)

- Only the transmitting or receiving antenna is moved along the measurement line while the other antenna stays stationary.
- In principle, a WARR measurement follows the same relationships concerning travel time as a CMP measurement. The difference is that the reflection point moves along the reflector. This is why a WARR measurement strictly is only applicable in the presence of horizontal or slightly sloping reflectors and material properties are

homogeneous.



 CMP and WARR measurements provide more information than a CO measurement. The drawback of these techniques is the high measurement effort since both procedures only provide point information for a specific location. Both method are hardly applicable along long measurement lines





Field example



Radargram Kaali Crater, Estonia

Energy loss and Penetration



- 1) Signal Attenuation (absorption)
- 2) Signal scattering
- 3) Spherical Spreading

Reynolds, (2011)

Energy loss and Penetration

• The electromagnetic signal is attenuated by different processes on its way through the soil.



Reynolds, (2011)

Energy loss and Penetration

• The largest amount of energy loss results from damping of free charge carrier movement. The material dependent attenuation is induced by the direct current electric conductivity σ_{dc} of the investigated medium. Depending on the traveled distance x the amplitude *E* of the electromagnetic wave decreases exponential with respect to its starting value E_0 :

$$E(x) = E_0 e^{-\beta x}$$
, with $\beta = \frac{\sigma_{dC}}{2c_0 \varepsilon_0 \sqrt{\varepsilon}}$

- The higher σ_{dc} of the medium, the higher is the attenuation of the electromagnetic wave. In soils, electrical conductivity for example increases due to an increasing in soil moisture content, clay content or amount of dissolved solutes in the solution.
- The penetration depth of the electromagnetic waves reduces with increasing electrical conductivity of the medium. For salt water the penetration depth is only 1cm:

$$\delta = \frac{1}{\beta}$$

Energy loss and Penetration



DEWOW



Pitfalls



100 MHz; time window: 400 ns; common offset profile; length: 59 m; DEWOW + Gain (attenuation = 4); v = 115 mm/s

Pitfalls



100 MHz; time window: 400 ns; common offset profile; length: 59 m; DEWOW + Gain (attenuation = 4); v = 115 mm/s











